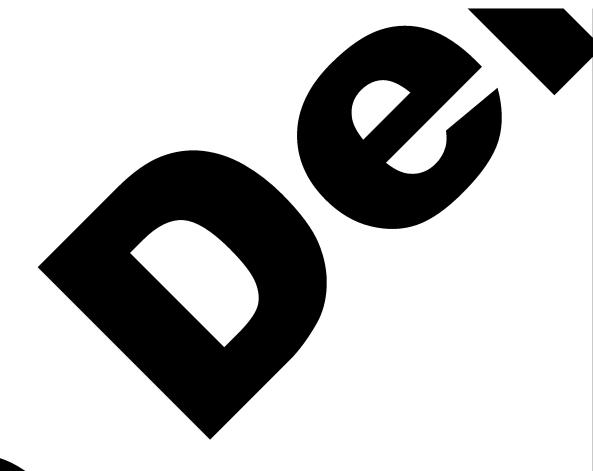
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CONTAMINATION OF OCEANS BY LONG-LEVED RADIONUCLIDES ACCORDING TO THE RESULTS OF SOVIET INVESTIGATIONS

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I. Introduction

Contamination by radioactive substances particularly by Sr^{90} , can inflict irreparable harm to the resources of the Globe Ocean (1).

The research conducted up to this day on radioactive contamination of oceans was chiefly aimed at appraising its hazards. On the other hand, radioactive "tags" in the ocean have to be used to obtain new data on the Processes occurring in it.

The present paper deals with the results of determination of Sr^{90} and Cs^{137} contents in the ocean water samples taken during the expeditions of the research ships "Vityaz" and "Mikhail Lomonosov" in 1959-1962.

The determination of Sr⁹⁰ was effected by Y⁹⁰ extracted from carbonate precipitate of sea water. The chemical yield of Sr was measured by the method of flame photometry. The content of Cs¹³⁷ in the sea water was determined by cobalt nitrite precipitation with the subsequent reprecipitation as bismuth icdide and chloroplatinate Cs and by extraction as ferrocyanide compounds (2,3).

The activity of the isolated samples was measured on counters with a background of about 1 count/min. (4,5) and on a low-background gamma-spectrometer with liquid scintillator (6).

Pacific Ocean

Tables I and II present the data on Sr^{90} content in the waters of the Pacific Ocean in 1961, i.e. after a two-year interval in nuclear tests. It is interesting to make a comparison between these results and the ones of early observations.

25 YEAR RE-REVIEW

In 1954 local sites of radioactive contamination appeared on the surface of the Ocean near the proving grounds.

The concentration of Sr^{90} in these areas attained values up to 5800 disintegrations/min. per 100 litres (7). For the last years reduction of the maximum values and general levelling of Sr^{90} concentration in the surface waters of the Pacific Ocean was ovserved (8).

In 1961, as seen from table I, the original contamination sites were found to be washed out due to horizontal and vertical intermixing of waters and motion of radioactive nuclides in the waters.

The maximum concentration of Sr⁹⁰ in the surface layer of the Ocean in I96I did not exceed I22 disintegrations/min per I00 litres and the mean value of the results for waters of the nothern hemisphere amounted to 50±IO disintegrations/min. per I00 litres.

Simultaneously with reduction of hetero geneity in Sr⁹⁰ distribution a further reduction of its mean concentration level in the surface layer in the western part of the Pacific Ocean is observed.

Along with this, in the eastern part of the Ocean some increase of concentration of ${\rm Cs}^{137}$ and consequently ${\rm Sr}^{90}$ was observed in the surface waters (IO). Hence, it appears that there was a considirable shifting of ${\rm Sr}^{90}$ with waters in the west-east direction.

A two-years period was not sufficient for making equal the concentration of Sr⁹⁰ in the northern and southern hemispheres. In the northern hemisphere waters in 1961 it was almost twice high than in the southern hemisphere.

The most interesting problem on vertical distribution of nuclear test products in the Pacific Ocean remains up to the present time poorly investigated.

The results given in table II allow to conclude that characteristics of the vertical distribution of Sr^{90} seem to differ in various points of the Ocean, and concentration decrease with depth occurs not always monotonously.

The speed at which radioactive contamination sirtes are washed away reduction of Sr⁹⁰ concentration in the surface waters of the Pacific Ocean indicate that water exchange between various parts of the Ocean in the northern hemisphere and between various depths does not take place as slowly as believed before.

Indian Ocean

The data on concentration of Sr^{90} in surface waters of the Indian Ocean are presented in tables III and IV.

The most typical feature of radioactive contamination of the Indian Ocean is the more or less uniform distribution of Sr 90 on the surface of the Ocean up to 40° latitude South. The resultant values of concentration fluctuate about 24±4 disintegrations/min per IOO litres. In view of the fact that the northern hemisphere accounts for but an insignificant part of the area of the Indian Ocean, the difference in the level of contamination of water in the north and the south, observed in other oceans, could not manifest itself here to the same extent.

The comparison of concentration of Sr^{90} in the surface waters with radioactive contamination of the air overwater layer in the north part of the Indian Ocean (II) indicated that mostly the increase of fragment content in the air leads to the increase of Sr^{90} concentration in the water. This correlation, however, is not absolute due to the complex mechanism of fragment migration in the Ocean, which assumes predomination of one process or the other depending on conditions.

It is well known that different density of radioactive fall-out along the meridian is observed on the surface of land even within one, in particular, the southern hemisphere (II). It may be supposed that the density of radioactive fall-out on the surface of the Indian Ocean also depends on the latitude. Then the data of table III testify to the existence of a considerable exchange of surface waters between various regions levelling of Sr⁹⁰ concentration. Biological factors as well as hydrological ones are of great importance in Sr⁹⁰ migration. Radiochemical analysis of corals and shell-fish collected in shoal-water in various regions of the northern part of the Indian Ocean showed a considerable increase of specific activity of Sr⁹⁰ in comparison with that (per I g. of stable Sr) in the surface waters; it indicates that there is a reservoir of Sr⁹⁰ conservation in the Ocean due biological precipitation of carbonates.

The specific activity of shell-fish was 98 disintegration/min/g of stable Sr (average from I2 samples), of corals - 61 disintegrations/min/g of stable Sr (average from I0 samples).

The determination of migration speed of Sr⁹⁰ measurements were made in the surface layer of globi gerina coze taken in the point I6°58' latitude north, 67° 30' longitude east at the depth 4000 m. The activity of Sr⁹⁰ was 5-6 disintegrations/min (I00 g of dry ground). This indicates a rather high speed of vertical migration of Sr⁹⁰ being just with biogenous character (skeletons of plankton foraminifera).

Proceeding from these data one can speak of a notable vertical transfer. The activity of Sr⁹⁰ is observed practically at all depths. The transfer coefficient is variable, the factors causing this transfer are of physicohydrological and biogenous character.

Atlantic Ocean

Table Y presents the results of determining the concentration of Sr⁹⁰ in the surface waters of the Atlantic Ocean for the period between September 1959 and December 1962.

However, in view of the fact that individual values in each ocean region do not differ greatly from one another, an averaging of the resulting data was made so as to trace the general nature of Sr^{90} change concentration in the surface layer.

For the sake of comparison, published data referring to I954-I958 were used (I2, I3, I4). The following pattern is given for the northern part of the Atlantic Ocean:

	I954 -	- 18 ·	disintegrations/min	per	100	1 (12)	single value
	I955 -	- II2	II			(I2)	_#_
	1956 -	- I2	_"_			(12, 13)	_#_
	195 7 -	- IO	"			(12, 14))
	1958 -	- 8	π17			(14)	
	1959 -	I 6	"n_		(the	preser	t paper)
	1960 -	21	_11_			-"-	
six	first months 1961 -	· 14	_11_			_"_	, i
six	last months	03	_"_				
	1961 - of 1962	-	_"_			_"_	
		'					

These results lead to the conclusion that despite the uninterrupted increase of deposited Sr^{90} , the concentration of Sr^{90} in the surface layer changes very alignly.

Of interest is the increase in the last six months of 1961 in connection with sharp rise in the intensity of the fall-out.

It may be presumed that the concentration of Sr 90 in the surface layer of water rises during the periods of an increased rate of fall-out and drops when the rate of fall-out diminishes.

There is no doubt that the absence of any substantial rise in the concentration of Sr^{90} in the surface waters testifies to the intensive removal of Sr^{90} from the surface layer of the Ocean to deeper layers.

In the waters of the tropical zone of the Atlantic Ocean, the concentration of Sr^{90} proved in 1961 and 1962 I,6 times lower than in the waters of the northern part.

An appreciable rise in concentration in I962 as compared with the first six months of I96I was exhibited there more vividly than in the northern part.

In the first six months of 1961, several determinations of the content of Cs¹³⁷ were made in the surface waters of the Atlantic Ocean. The results are given in Table YI. All the explored points, with the exception of one, belong to the tropical part of the Ocean. The average value for the points amounted to 20 disintegrations/min per 100 litres. In the same zone and for the same time period, the concentration of Sr⁹⁰ (Table Y) amounted to 9 disintegrations/min per 100 litres. The ratio of these figures (2.2) quite approximates that of the activities of Cs¹³⁷ and Sr⁹⁰, peculiar of fission products.

The results of measurements of Sr^{90} concentration of various depths are given in Table YII. The surface layer down to 75-I00 m generally has the same concentration of Sr^{90} in all horizons as in the Indian Ocean. With a greater depth, the concentration, as a rule, diminishes but not always uniformly.

This may occur where waters at a depth may have come down from the areas of a greater contamination. Such a drop in the concentration of Sr⁹⁰ in proportion to depth attests that the spreading of this radionuclide to the depth is due not only to turbulent diffusion, but also to other processes.

The results given in the last column of Table VII indicate that in a layer of water down to depths of 500 to 1,000 m, the reserve of $\rm Sr^{90}$ at the explored points of the Ocean are rather great, for exceeding those deposited on land. This phenomenon should be taken into account when estimating the world reserve inventory of $\rm Sr^{90}$ (15).

Along with the open part of the Atlantic Ocean, samples of water were also taken in its closed seas: the Baltic Sea, the North Sea, the English Channel, the Mediterranean Sea, the Sea of Marmora and the Black Sea. It has been established that the concentration of Sr⁵⁰ in these seas is almost double that in the open ocean at the same latitudes. Evidently this is accounted for by the smaller volume of the water mass accessible for the intermixing of Sr⁹⁰ fallen out: the seas have no intensive exchange with less contaminated tropical water masses.

Conclusion

The results presented in this paper of determining radioactive contamination or the oceans cover rather great regions of the water space of each of them. Nevertheless, the number of explored points for these vast spaces should be considered quite insignificant. Publication of additional works on studies of radioactive contamination of oceans will apparently permit to present a more deteatled picture of radionuclide distribution.

From the data on the concentration of Sr⁹⁰ and its changes a notable speed of exchange is observed. Vertical distribution of this radionuclide indicates that it is difficult to regard deep waters as being isolated from the surface.

The research of the years in future includes the task of mastering the mechanism of horizontal and vertical transfer of artificial radionuclides in oceans.

But it is already clear now that the high speed of exchange raises doubts regarding the possible burial of waste of the atomic industry and transport in oceans.

Table I Concentration of Sr⁹⁰ in the surface waters of the Pacific Ocean in I96I

	Coordin	ates	Sr ⁹⁰ act-	•	Coordi	Sr ⁹⁰ act-	
Data	Latitude	Longitude	ivity dis/min per IOO litres	Data	Latitude	Longitude	ivity dis/min per IOO litres
I96I 07.04	N 04 ⁰ 13,	E I28 ⁰ 59•	43 <u>+</u> 7	02.II	12 ⁰ 51,	176 ^E 06•	22 <u>+</u> 6
	10 ⁰ 24,	130 ⁰ 23•	38±4		II ⁰ I5,	_	
09.04 I0.04	10°24', 14°00',	130°23'	38 ∸ 4 35 <u>+</u> 9	03.II 04.II	09 ⁰ 4I,	176 ⁰ 16, 176 ⁰ 03,	63 <u>+</u> 1I
II.04	17 ⁰ 50,	130 52,	42 <u>+</u> 9	04.II	07057	176°03' 175°57'	32 <u>+</u> 9 59+I2
I2.04	2I ⁰ I3,	130 ⁰ 45 •	124 <u>+</u> 33	05.II	05058,	175 ⁰ 04	44±9
I3.04	25°3I,	I30°58,	62 <u>+</u> II	07.II	030581	176 ⁰ 08,	53 <u>+</u> I0
I4.04	29033,	131°00,	40 <u>+</u> 7	07.II	0I ⁰ 57,	176 ⁰ 06,	3I <u>+</u> 5
I5. 08	34 ⁰ 54	I65 ⁰ 27,	2I <u>+</u> I0	08.11	00050	I76 ⁰ 02,	86 <u>+</u> 7
03.09 II.09	N 20 ⁰ 59' I3 ⁰ 58'	w I58 ⁰ II, I40 ⁰ 25,	57 [±] II 79 <u>+</u> I4	09.II I0.II II.II	s 00 ⁰ 03; 00 ⁰ 20; 03 ⁰ 58;	w 176 ⁰ 07 • 176 ⁰ 07 • 176 ⁰ 04 •	22 <u>+</u> 6 18 [±] 4 38 <u>+</u> 12
12.09	12 ⁰ 07,	I40 ⁰ 38'	32 <u>+</u> 6	I3.II	08°01'	175°57	30 <u>+</u> 7
I4.09	10 ⁰ 29'	I40 ⁰ 00'	26 <u>+</u> 4	I4.II	10 ⁰ 04,	176 ⁰ 00'	37 <u>+</u> 8
16.09	05 ⁰ 57	139 ⁰ 57,	20 <u>+</u> 4	14.1I	12 ⁰ 07,	175 ⁰ 59	2I <u>+</u> 5
21.09	00°00°	I40 ⁰ 02'	64 <u>+</u> I0	16.11	14 ⁰ 00'	176 ⁰ 08,	30 <u>+</u> 8
	S	W			S	E	
25.09	05 ⁰ 02	139 ⁰ 54,	43 <u>+</u> 8		•	•	
0I.IO	17 ⁰ 50	I40 ⁰ I7,	24 <u>+</u> 4	23.II	II ⁰ 58'	170°20°	15 <u>+</u> 3
03.10	18 ⁰ 04,	I48 ⁰ 43'	2I <u>+</u> 6	24.II	08 ⁰ 46'	166 ⁰ 41	36 <u>+</u> I3
15.10	13 ⁰ 02, 09 ⁰ 58,	154 ⁰ 01,	34 <u>+</u> 5	26. ₁ I	03 ⁰ 461	165 ₀ 01,	5I <u>+</u> I3
I6.I0	09°58°	154 ⁰ 02°	I5 <u>+</u> 4	00 T.	N 000.01	E	# O . O
18.I0 19.I0	05 ⁰ 03,	I54 ⁰ 02 ' I54 º II,	37 <u>+</u> II 20 <u>+</u> 5	27.I1 28.II	00°22,	162 ⁰ 30' 162 ⁰ 43'	49 <u>+</u> 8
	05 05 00 00 00 00 00 00 00 00 00 00 00 0	154 ⁰ 091	54 <u>+</u> 17	0I.I2	06 ⁰ 33,	162 43, 161 ⁰ 57,	30 <u>±</u> 5 94 <u>+</u> I5
24.IO	06 ⁰ 26	153 ⁰ 49'	53 <u>+</u> I1	02.12	07 ⁰ 39'	161 57	86 <u>+</u> 14
	09 ⁰ 46,	153 ⁰ 54'	54 <u>+</u> I0	07.12	16 ⁰ 25,	160°01,	87 <u>+</u> 15
	• •	384	- 7				~ 1 2 2 2

29.10	12 ⁰ 43,	159 ⁰ 03•	29 <u>+</u> 6			155°35°	
	13 ⁰ 02,	165°21'	27 <u>±</u> 6		_	151 ⁰ 53,	
31.10	I3 ⁰ 05•	170 ⁰ 2I,	33 ±7	15.15	34 ⁰ I0,	I4I ⁰ 27•	20711

Table II

Concentration of Sr⁹⁰ in deep waters of the Pacific Ocean, I96I

Data	Coordinates			Horizon	Sr ⁹⁰ activity	
					mm	disintegrations/ min per IOO 1
9.04.1961	10 ⁰ 24'	n.lat	130°23'	E.long	0	38 ± 4
					900	I,4 ± I,0
					3000	14 ± 4
12.04.1961	210131	n.lat	I30 ^o 45 '	E.long	0 900 5500	124 ± 33 9 ± 1 4 ± 1

							
			Sr ⁹⁰ act-				Sr ⁹⁰ act-
Le	្ 	Cocrdinates	ivity dis/min per IOO litres	Data	Coordi	nates	ivity dis/min per IOO litres
1959	S.lat	E.long			n.lat	E.long.	•
3.II	8 ⁰ 00,	109 ^C 40'	30 <u>+</u> 2	25 . I0	I4 ⁰ 44,	50 ⁰ 02 °	20 <u>+</u> 5
5.I1	10°20°	10 ⁰ 40°	36 <u>+</u> 3	28.10	II _O II,	52 ⁰ I3,	25 <u>+</u> 8
25.II	7004,	99 ⁰ 42•	1 <u>+</u> 81	29.10	10 ₀ 19,	53 ⁰ I4 '	I8 <u>+</u> 6
29.II	16 ⁰ 02,	99 ⁰ 48•	∠2 <u>+</u> I	31.10	9 ⁰ 23	54 ⁰ 56°	2I <u>+</u> 4
30.II	15 ⁰ 52,	104 ⁰ 50°	38 <u>+</u> 2	I.II	9 ⁰ 33,	57 ⁰ 08 ,	41 <u>+</u> IO
2.12	23 ⁰ 39•	113041,	28 <u>+</u> 2	3.II	I5 ⁰ 23	58 ⁰ 23 '	49 <u>+</u> I0
I7.12	30°2I,	89 ⁰ 53•	I5 <u>+</u> I	5.II	22 ⁰ 23•	59 ⁰ 53 ,	12 <u>+</u> 3
21.12	Ιο ⁰ 09,	90°18°	36 <u>+</u> 3	8.11	22 ⁰ 49	62 ⁰ 03;	I4 <u>+</u> 2
				IO.II	15 ⁰ 03	62°00'	I3 <u>+</u> 3
1960	n.lat.	E.long					
7.0I	5 ⁰ 15,	84 ⁰ 57•	I8 <u>+</u> 2	13.11	8 ⁰ 40°	61 ⁰ 54°	32 <u>+</u> 5
	S.lat.	E.long.		15.II	3°25°	62 ⁰ 06°	23 <u>+</u> 5
I0.0T	0022	85 ⁰ 19•	23 <u>+</u> 2		S.lat	E.long.	
I5.01	I6 ⁰ 58,	67 ⁰ 30°	20 <u>+</u> 2	17.TI	0 ⁰ 55•	62 ⁰ 33,	2I <u>+</u> 5
I9.0I	8 ⁰ 27,	75 ⁰ 50 °	I5 <u>+</u> I		n.lat	E.long	-
	n.lat.	E.long.	_	21.11	6 ⁰ 59,	65°57°	28 <u>+</u> 4
27.0I	6 ⁰ 52 ,	76 ⁰ I2 °	6 <u>+</u> I	26.II	19 ⁰ 15,	65 ⁰ 56 °	20 <u>+</u> 5
4.02	10 ⁰ 46,	75 ⁰ 48•	5 <u>+</u> I	12.12	0°00°	70 ⁰ 52°	24 <u>+</u> 5
					S.lat.	E.long	;
1 960	S.lat	E.long.		I4.I2	3 ⁰ 56,	73 ⁰ 10,	4 I± 9
12.02	3 ⁰ 02,	65 ⁰ 02 '	20 <u>+</u> I	18.12	10 ⁰ ,40°	70 ⁰ 59°	2I <u>+</u> 7
16.02	8 ⁰ 52	67 ⁰ 30°	19 <u>ŧ</u> I	20.12	I6 ⁰ 42,	70 ⁰ 58°	36 <u>+</u> 7
I9.02	15 ⁰ 56'	67 ⁰ 23 '	17 <u>+</u> I	22.12	24 ⁰ 04 ,	71 ⁰ 17,	3I <u>+</u> 7
22.02	16°10,	52 ⁰ 42 ,	6 <u>+</u> 2	24.I2	30 ⁰ 00 •	7I ⁰ 22,	I8 <u>+</u> 9
2.03	I4 ⁰ 5I '	53 ⁰ 0I,	17 <u>+</u> I	26.12	36 ⁰ 06,	7I ⁰ I8,	8 <u>±</u> 3
	n.lat.	E.long.		27.12	39 ⁰ 24 ¹	7I ⁰ 19,	19 <u>+</u> 4
26.03	5 ⁰ 04,	58 ⁰ 0 7 •	38 <u>±</u> 3	29.12	3I ⁰ 50,	80 ⁰ 35 •	23 <u>+</u> 6
2.04	I4 ⁰ 03•	68 ⁰ I7• ³84	44 <u>+</u> 2	31.12	24 ⁰ 44 •	83 ⁰ 03•	I0 <u>+</u> 5

I3.04	16 ⁰ 25'	61 ⁰ 51,	20±3				
1961 10.01	S.lat I ⁰ 57, n.lat	E.long 83 ⁰ 03, E.long.	3ა <u>+</u> 5	28.02	n.lat. 18 ⁰ 03, 9 ⁰ 14,	E.long. 90 ⁰ 54, 93 ⁰ 40,	I3±8
24.0I	4 ⁰ 22'	83 ⁰ 05,	I4 <u>+</u> 4	6.03 IO. 03	9°14' 0°56'	93°40° 92°35°	56 <u>+</u> 11
28.0I	13 ⁰ 17,	83 ⁰ 03•	40 <u>+</u> 9	10.00	0 50	<i>y</i> 2 3 <i>y</i> ,	46 <u>+</u> 9
3I.OI	18 ⁰ 58'	87 ⁰ 00•	28 <u>+</u> 6	13.03	ol _o oo,	96 ⁰ 56•	33 <u>+</u> 6
2.02	13 ⁰ 37,	86 ⁰ 53 ,	6 <u>+</u> 3	I7.03	5 ⁰ I7'	100 ⁰ 37,	23 <u>+</u> I0
5.02	6 ⁰ 62'	87 ⁰ 02,	24 <u>+</u> 4	20.03	I3 ⁰ 40,	I04 ⁰ 49,	34+10
I4.02	II ⁰ 04'	92 ⁰ 02•	20 <u>±</u> 6	22.03	8 ⁰ 10,	I04 ⁰ 39,	I9+3

Concentration of Sr^{90} in the deep waters of the Indian Grean

Data	Coordinates		morizon	Sr ⁹⁰ activity disintegrations/min per IOO litres
3.11.59	E°(('S.lat.	109 ⁰ 40' E.long	• 0	30 ± 2
			300	25 ± 2
5.41.59	IO ⁰ 20' S.lat.	107°40' E.lon	g. 0	36 ± 3
			300C	2 ± I
			6200	3 ± I
17 - 12.59	30°21' S.lat	89°53' E.long	0	15 ± 1
			100	20 ± 3
			1000	4 ± I
21-12.59	16 ⁰ 09' S.lat	90°18' E.long	0	36 ± I
			100	17 ± 1
			1000	10 ± 1
	_	_	4000	3 ± I
10.01.60	O ⁰ 22' S.lat	85°19' E.long	0	23 ± 2
			100	24 ± I
			1000	IO ‡ I
19.01. 60	8°27' S.lat	75°50' E.long	0	15 ± I
			47	16 ± 1
			200	9 ± 2
	•	_	1000	4 ± I
27. OI.60	6°52' n.lat	76°12' E.long	0	6 ± I
			200	6 ± I
	•	_	1860	9 ± I
12-02-60	3'02' n.lat	65°20' E.long	0	20 ± I
	•	_	3500	4 ± I
19.02.60	15°56' S.lat	67°23' E.long	0	17 8 I
			I50	2I - I
2.04.60	I4 ⁰ 03' n.lat	68°17' E.long	0	44 ± 2
	384		200	58 ± 3

Concentration of Sr⁹⁰ in the surface waters of the Atlantic Ocean, 1959-1962

n=+ -	Coordin	ates	Sr ⁹⁰ act-		Coordina	ites	Sr ⁹⁰ act-	
Data	Latitude	Longitude	ivity dis/min per IOO litres	Date	Latitude	Longitude	ivity dis/min per IOO litres	
1959	North atla	ntics		1961	n	W		
	n	w		17.03	_	10 ⁰ 34 '	II	
16.09	36 ⁰ 50'	58 ⁰ 281	IO	19.03	_	16 ⁰ 58'	II	
21.09	36 ⁰ 52'	68 ⁰ 12'	23	17.06		30°03'	21	
		···		17.06	_	30°00'	II	
avere	ge according vations in	ng to 1959	16	1962	n n	w	4.4	
				26.12	_ `	08 ⁰ 591	25	
1960	n	₩						
22.OI	45 ⁰ 48 '	13 ⁰ 56 '	60	Averag	ge accordin	g to		
23.01	43°02'	21 ⁰ 04'	9	observ	vations at	the end		
24.OI	_	37 ⁰ 491	9	of 196	52	••••••	24	
31.01	33°07 '	50°36'	18					
03.02	_	62 ⁰ 18'	23	1961	n	**		
03.02	28 ⁰ 39 '	62°18'	28	20.06		29 ⁰ 57	20	
03.02	28 ⁰ 39 1	62 ⁰ 18'	28	20.06	ن 57°39	39 ⁰ 57 '	17	
10.02	27°03 •	66°25 '	9	22.06	43°13'	2 1°26 •	9	
13.02	_	76°20 '	7	25.06	50°10'	02 ⁰ 26 •	12	
16.02	31 ⁰ 57 '	78°18'	IO	Averso	e accordin	r +0		
21.02	38 ⁰ 59 '	71 ⁰ 59 '	53		ations of	-		
23.02	37°32 •	70 ⁰ 15 '	IO		nths of 19		Τ./	
04.04	54 ⁰ 52 •	13 ⁰ 51'	12		_		14	
				27.09		36°031	14	
		to observat	ions	27.09		36°03 •	40	
in I	960	• • • • • • • • • • • • • • • • • • • •	. 21		37°33'	58°37'	18	
					37°33'	58 ° 37 '	17	
					37°51'	58°07'	19	
				-	37°51'	58 ⁰ 09'	IG	
		384			37°51'	58°07'	15	
					39° 291	62 ⁰ 55'	25	
				2I.09	390291	62 ⁰ 55 •	30	

				22.09	39 ⁰ 29+	62 ⁰ 55 1	29
				25.09	430201	520001	25
				25.09	450501	51° 30 •	T9
				07.10	44°20'	52°11'	26
				0,020	,,	,	
				Avera	ge accordi	ng to	
					vations of		
				six m	onths of I	961	23
1962	n	W					
20.10	38 ⁰ 221	10 ° 16 •	24	29.05	I6 ⁰ 27'	30°01'	G
25.IO	270221	14°07'	24	30.05		30 ⁰ 00+	G
24.I2	25° 21'	15 ⁰ 48 •	18	31.65		30°03 •	6
25.12	29°04'	12 ⁰ 39 '	27	03.06	_	30°08'	8
196 I	n	V		04.06	_	30°06'	4
09.06	_	30°06 •	6	04.06		30°00'	7
09.06		30°06 '	II	05,.06	_	30°00'	18
12.06		29 ⁰ 56 '	16	1962			
12.06		29 ⁰ 51'	8	13.12	s 16 ⁰ 02'	W 24 ^Q 224	=-
15.06		30°00 •	IO	-		34 ⁰ 00'	10
15.06	_	30°03 •	II	13.12	16 02'	34°00'	9
			· · · · · · · · · · · · · · · · · · ·	_	n	₩	
-	e according			16.12		27°27'	13
		he first six		12.12		25° 27 '	IO
	of 1961 Quatorial a	tlentice	9	22.12	-	20°361	31
1962	n n	W CIGHTER		23.12	20 ⁰ 52'	18 ⁰ 40'	9
22.03	16 ⁰ 58'	20 ⁰ 29 '	15	Aver	age accord	ing to	
27.03	II°00'	18 ⁰ 15'	6		rvations a	•	
04.04	ου ^ο 17 '	20 ⁰ 19 •	13		962		16
07.04	00°03'	15 ⁰ 04 •	16	1962	n	W	
	S	W		26.10		17 ⁰ 27 '	17
II.04	00°2I'	04 ⁰ 26 •	IO	2 6. I		18 ⁰ 14'	16
II.04	00°2I'	04 ⁰ 26 •	4	31.10		17 ⁰ 55'	21
02.05	10003'	20°33 •	6	01.1	_	18°10'	II
02.05	10003'	20°33 •	6	08.1		18010	27
05.05		27013 •	13	 -	5	₩	
Q6.Q5	20001'	30°03'	8	II.II	_	25°10'	33
07.05	200011	30 ⁰ 03 *	5	17.11	_	30°18'	17
07-05	20001	30 ⁰ 03 •	7	20.11	_	34 ⁰ 55 '	16
29.05	16 ⁰ 25'	30°03 '	II	20.11	_	34 ⁰ 55 •	16
		384					~~

13	38 ⁰ 05 '	10 ⁰ 39 •	21.II
16	31 ⁰ 50'	12°15 •	22.II
IO	34 ⁰ 54'	12°04 •	23.11
16	33°2I'	16 ⁰ 00 •	28.II
13	39°51'	22 ⁰ 14 •	II.12
19	39 ⁰ 51'	22°14'	11.12
	nti c s	uth atlen	So
	W	S	1961
IO	29 ⁰ 541	23°03 •	II.05
16	30 ⁰ 05	29 ⁰ 58 •	12.05
4	30°17'	29 ⁰ 51 '	12.05
9	30°17'	29 ⁰ 51 •	12.05
7	40°00'	25° 17 •	I5. 05
13	39°58'	25°18•	16.05
IO		age	Aver

Table VI

Concentration of Ca^{I37} in the surface waters of the Atlantic Ocean

Data	Coordinates	Cs 137 activity disintegrations/min per 100 1
19.03.1961	24°30' nelat 16°57' welong	30
22.03.1961	20 ⁰ 27'"" 17 ⁰ 50'""	25
7.04.1961	00°02' S.lat. 13°41' " "	24
7.05.1961	20°0 1 '"" 30°07'""	17
31.05.1961	09 ⁰ 46'"" 30 ⁰ 02'""	8
II.06.196I	I0 ⁰ 59' S.lat. 29 ⁰ 57' " "	20
15. 06.1961	2I°24' " " 30°03' " "	22
	average	21

Data	Coordi	inates	Horizon m	Sr ⁹⁰ act- ivity dis/min per IOO litres	Stooks of Sr ⁹⁰ in the explored layer, micro- curies/sq.km
4.04.61	00°17'n.lat 2	20°19' w.long	0	13	
5.04.61			30	16	6.2
4.04.61			50	17	
4.04.61			100	7	
7.05.61	20°01' S.lat	30 ⁰ 03' w.long	0	5	
			0	7	3.2
			30	8	
			50	7	
			100	7	
12.06.61	II ⁰ 0I' n.lat	29°56' w.long	0	16	
			30	9	4.4
			100	7	·
I6. I0.62	60°12' n.lat	OI ⁰ 39' w.long	Ο '	38	
20.10.62	38 ⁰ 22' n.lat	IOOI6' w.long	0		26
I8.10.62	44°40' n.lat	9°17' w.long	70	24	
18. 10.62	44°40' n.lat		200	31	
20.10.62	38°22' n.lat	IOOI6' w.long	O	24	
			50 0	9	64
			1000	13	
26.10.61	20 ⁰ 29' n.lat	17°27' w.long	O	17	
31.10.62	10000, n.lat	I7055, w.long	25	I5 II	3 . 3
31.10.62	IO 00 n.lat	17055' w.long	50 0	11 2I	
71,10,00	10 00 11.140	17 // "1016	50 200	II 6	9•4
08.11.62	OI ⁰ 05' n.lat	I8º10' w.long	O	27	
06.11.62	OI°00' S.lat	18000' w.long	50	12	
07.11.62	00 ⁰ 05' n.lat	18°16' w.long	85	29	16
			190	II	
	384		200	II	

II.II.62 O	I ⁰ 00' S.lat 25 ⁰ IO' w.long O	33
	100	14 36
	500	14
28.II.62 I	6°00' S.lat. 33°21' w.long 0	17
	100	6 32
	800	6
17.12.62 00	⁰ 52' n.lat 25 ⁰ 27' w.long 0	IO
	50	10 13
	300	8

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